



## Pedogenic iron oxides in soils of the Acre State, Brazil

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**ABSTRACT:** Iron oxides are environmental indicators and influence on physical and chemical behavior of soils. This research aimed to identify and characterize pedogenic iron oxides in soils of the Acre state, Brazil. The soils developed from sedimentary rocks of the Solimões Formation. Twenty one samples of subsurface horizons were collected from ten soil profiles. Soil samples were analyzed by x-ray diffractometry associated to selective chemical dissolutions. Fe contents in the clay minerals (Fes), in the pedogenic iron oxides (Fed) and in the poorly crystalline oxides (Feo) were determined. The Al<sup>3+</sup>-substitution and the mean crystal diameter of goethite and hematite were estimated. Soils presented low contents of Fes and Fed. The Fed/Fes ratio indicated soils in the intermediate stage of weathering, with dominance of goethite, hematite and maghemite. This condition of weathering was confirmed by the higher frequency of goethites and hematite with intermediate Al<sup>3+</sup>-substitution. Goethites have isodimensional forms and hematite have flat plaques forms.

**Key words:** mineralogy, x ray diffraction, hematite, goethite.

## Óxidos de ferro pedogênicos em solos do estado do Acre, Brasil

**RESUMO:** Os óxidos de ferro são indicadores ambientais e influenciam o comportamento físico e químico dos solos. Este trabalho objetivou identificar e caracterizar os óxidos de ferro pedogênicos de solos do estado do Acre, desenvolvidos de rochas sedimentares da Formação Solimões. Vinte e uma amostras de horizontes subsuperficiais de dez perfis de solos foram analisadas por difratometria de raios x associada a dissoluções químicas seletivas. Os teores de Fe total (Fes), dos óxidos de Fe pedogênicos (Fed) e de baixa cristalinidade (Feo) foram determinados; e a Al<sup>3+</sup>-substituição e o diâmetro médio cristalino da goethita e da hematita foram estimados. Em geral, os solos apresentaram baixos teores de Fes e de Fed. A relação Fed/Fes indicou solos em estágio de intemperização intermediário, com dominância dos óxidos de Fe goethita, hematita e maghemita. Essa condição de intemperismo foi reforçada pela maior frequência de goethitas e hematitas com intermediária Al<sup>3+</sup>-substituição. As goethitas possuem formas isodimensionais e as hematitas ocorrem como placas achatadas.

**Palavras-chave:** mineralogia, difratometria de raios x, hematita, goethita.

Formation and stability of Fe oxides are governed by specific pedoenvironmental conditions, which makes the oxides useful indicators for pedoenvironmental and pedogenetic processes (KÄMPF & CURI, 2000). Usefulness of iron oxides arises from their widespread occurrence in various forms and concentrations, and their degrees of crystallinity and isomorphic substitution of Fe by other cations (INDA JUNIOR & KÄMPF, 2005). Also, Fe oxides have physical and chemical effects on soil that depend on their specific surface area, reactivity and variability in surface charge.

Studies on Fe oxides have facilitated the interpretation of past and present pedogenetic processes (INDA JUNIOR & KÄMPF, 2005), and

the mapping and classification of soils (SANTOS et al., 2013). Whereas hematite, goethite and maghemite prevail in aerobic pedoenvironments, magnetite, which is a mineral of lithogenic origin, usually occurs at lower concentrations and ferrihydrite is typically present in small amounts. In cyclically anaerobic pedoenvironments, goethite, lepidocrocite and ferrihydrite are abundant, whereas hematite is restricted to mottles and nodules. Other mineralogical properties such as crystallinity and the degree of Al<sup>3+</sup> substitution (viz., the extent to which Fe is replaced with Al) are also effective environmental indicators and reflect specific mineral formation conditions. These mineral properties can be assessed by X-ray (XRD) diffraction spectroscopy,

which allows mean coherence lengths (MCL) and  $Al^{3+}$  substitution to be estimated (SCHULZE, 1984; SCHWERTMANN et al., 1979).

In this research, we used the XRD technique in combination with selective dissolution treatments to identify and characterize pedogenic Fe oxides in soils from the Brazilian state of Acre. Twenty one samples were collected from the subsurface horizons of soil profiles located between the municipality of Cruzeiro do Sul and the city of Rio Branco along the BR 364 road (ANJOS et al., 2013). Soils originated from Cenozoic sedimentary rocks including lacustrine argillites, siltites and sandstonites in the Solimões formation under the influence of sedimentary deposits in the foothills of the Andean mountain range. The target soils were examined at the IX Brazilian Meeting for Soil Classification and Correlation in 2010.

Fe forms were selectively extracted from the air-dried fine earth fraction (ADFE, particle diameter <2mm). The amount of Fe extracted by sulfuric acid (DONAGEMA et al., 2018) was taken to represent total Fe (Fes). The combined amount of pedogenic oxides (Fed) was extracted with sodium dithionite–citrate–bicarbonate (DCB Na) at 80°C (MEHRA & JACKSON, 1960), and that of low-crystallinity oxides (Feo) with 0.2mol L<sup>-1</sup> ammonium oxalate at pH3 in the dark (MCKEAGUE & DAY, 1966). Finally, all dissolved Fe forms were determined by atomic absorption spectroscopy (AAS), and total organic carbon (TOC) in ADFE was quantified by oxidation with potassium dichromate (ANJOS et al., 2013).

The Fe oxide fraction was concentrated by addition of a hot 5mol L<sup>-1</sup> solution of NaOH (KÄMPF & SCHWERTMANN, 1982). XRD spectra were acquired with a Philips diffractometer equipped with a Fe monochromator, using Fe-K $\alpha$  radiation at 30kV at 30mA. Measurements were made on non-oriented slides of the concentrated iron oxide fraction over the  $\theta$ 20 range 25–50. For this purpose, an amount of 300mg of sample was mixed with 30mg of halite (NaCl) that was used as internal standard.

The hematite/(hematite+ goethite) ratio, Hm/(Hm + Gt), was estimated from the area ratio of the line Hm 104 or Hm 012 to the line Gt 100, and the degree of  $Al^{3+}$  substitution in hematite was calculated from the d(110) spacing (SCHWERTMANN et al., 1979) in hematite, and the d(110) and d(111) spacing in goethite (SCHULZE, 1984). Finally, crystallinity was estimated from the mean coherence lengths (MCL) in the hkl planes

110 and 111 in Gt, and 104 and 110 in Hm (INDA JUNIOR & KÄMPF, 2005).

As can be seen from Table 1, Fes values ranged from 6.3g kg<sup>-1</sup> in profile AC 01 (sandy sediments) to 73.4g kg<sup>-1</sup> in AC 06 (sandstone and clay sediments). Based on the Brazilian System of Soil Classification (SANTOS et al., 2013), the Fes values were typical of mesoferric soils (80g kg<sup>-1</sup><Fe<sub>2</sub>O<sub>3</sub><180g kg<sup>-1</sup>) in five horizons and hypoferric soils (Fe<sub>2</sub>O<sub>3</sub><80g kg<sup>-1</sup>) in all others. Fed values ranged from 0.6 to 43.3g kg<sup>-1</sup> and Feo values from 0.2 to 2.7g kg<sup>-1</sup>. The Fed/Fes ratio spanned the range 0.10–0.67 but most values were lower than 0.5 (Table 1), which suggests that a substantial proportion of Fe was present as primary and secondary aluminosilicates (KÄMPF & CURI, 2000). Based on the negative correlation with index  $K_i$  ( $R=-0.791$ ,  $P<0.001$ ,  $n=15$ ) previously reported by ANJOS et al. (2013), the Fed/Fes ratio was a useful indicator of weathering in most of the studied soils. The Feo/Fed ratio range from 0.01 to 0.83 but was less than 0.10 in most cases, suggesting a prevalence of crystalline forms (goethite, hematite and maghemite); also, the contents in poorly crystalline oxides —ferrihydrite, mainly— were comparatively low (SCHWERTMANN et al., 1982). The prevalence of poorly crystalline oxides in profile AC01 was a result of the restrictive effect of organic matter on iron oxide crystallinity as suggested by the positive correlation between the Feo/Fed ratio and the TOC content:  $Feo/Fed=-0,158+0.0656 \cdot TOC$ ,  $R=0.861$ ,  $P<0.001$ ,  $n=16$ .

The XRD spectra, not shown, allowed the Fe oxides goethite, hematite and maghemite to be identified. The Hm/(Hm + Gt) ratio ranged from near zero in horizons Biv and BCv2 of profile AC 07 to 0.60 in horizon Bt2 of AC 10 (Table 2). The prevalence of goethite over hematite in most of the samples was a result of the low Fe contents of the original materials (KÄMPF & CURI, 2000). Sample color, which spanned a wide range of hues (2.5–10 YR), was positively correlated with the Hm/(Hm + Gt) ratio:  $Hue=10.736-[14.809 \cdot Hm/(Hm + Gt)]$ ,  $R=0.773$ ,  $P<0.001$ ,  $n=17$ ). Most of the samples contained maghemite, the main pedogenic mineral with ferrimagnetic properties in the clay fraction.

As can be seen from Table 2, the estimated degree of  $Al^{3+}$  substitution in goethite differed widely among samples. The most frequent values fell in ranges suggesting (a) the formation of crystals in redoximorphic pedoenvironments or in soils from materials containing little  $Al^{3+}$

(Al<sup>3+</sup> substitution <15mol mol<sup>-1</sup>); and (b) the presence of moderately developed soils (Al<sup>3+</sup> substitution =15–20mol mol<sup>-1</sup>) (FITZPATRICK & SCHWERTMANN, 1982). Only samples from profiles AC 02 and AC 11 exhibited Al<sup>3+</sup> substitution values above 20mol mol<sup>-1</sup> and hence suggestive of highly weathered soils. The Al<sup>3+</sup> substitution values for hematite, 0.05–0.08mol mol<sup>-1</sup>, fell in the central range. The goethite-to-hematite Al<sup>3+</sup> substitution ratio exceeded 0.5 in most samples, which suggests that the two minerals formed at different times under also different pedoenvironmental conditions (INDA JUNIOR & KÄMPF, 2005).

Crystallinity was estimated from the mean coherence lengths (MCL<sub>hkl</sub>, Table 2). MCL<sub>110</sub> and MCL<sub>111</sub> for goethite ranged from 27 to 74nm and 86 to 158nm, respectively. In those samples where both parameters could be determined, MCL<sub>111</sub> exceeded MCL<sub>110</sub>, which is consistent with isodimensional development in the three crystallographic directions (SCHWERTMANN & KÄMPF, 1985). The MCL<sub>104</sub> and MCL<sub>110</sub> values for hematite spanned the ranges 69–154 and 29–59nm, respectively. The fact that MCL<sub>104</sub> values exceeded MCL<sub>110</sub> values suggested that hematite crystals were longer than they were high and hence had the typical appearance of flat

Table 1 - Hue, total organic carbon (TOC), Fe contents in selectively extracted Fe oxides forms and their ratios, and index  $K_i$ .

Profile	Soil	Horizon	Hue	-----Air-dried fine earth-----						$K_i$
				TOC	Fes (Fe <sub>2</sub> O <sub>3</sub> )	Fed	Feo	Fed/Fes	Feo/Fed	
-----g kg <sup>-1</sup> -----										
AC 01	EKo	Bhsx	10YR	11.8	6.3 (9.0)	0.6	0.5	0.10	0.83	1.68
AC 02	LAd	Bw2	10YR	1.2	21.0 (30.0)	12.1	0.2	0.58	0.02	1.70
AC 04	PVal-1	Bt1	2.5YR	3.9	42.0 (60.1)	25.5	1.2	0.61	0.05	1.78
AC 04	PVal-1	Bt2	2.5YR	3.7	64.3 (91.9)	43.3	0.5	0.67	0.01	1.43
AC 04	PVal-1	BCf1	2.5YR	2.4	72.7 (104.0)	40.7	0.7	0.56	0.02	2.04
AC 05	PVal-2	BA	5YR	4.4	33.6 (48.1)	14.3	1.6	0.43	0.11	3.42
AC 05	PVal-2	Bt1	2.5YR	3.6	42.7 (61.1)	20.0	1.0	0.47	0.05	3.16
AC 05	PVal-2	Bt2	2.5YR	2.9	54.5 (77.9)	–	–	–	–	3.37
AC 06	PVal-3	BA	2.5YR	7.2	40.6 (58.1)	–	–	–	–	2.89
AC 06	PVal-3	Bt1	2.5YR	5.9	47.6 (68.1)	–	–	–	–	2.71
AC 06	PVal-3	Bt3	2.5YR	3.3	73.4 (105.0)	32.9	0.9	0.45	0.03	3.25
AC 07	VXo-1	Biv	10YR	3.5	49.7 (71.1)	11.8	1.3	0.24	0.11	–
AC 07	VXo-1	BCv2	10YR	1.2	46.2 (66.1)	–	–	–	–	–
AC 08	TCp	Bt1	10YR	4.3	43.4 (62.1)	16.8	1.5	0.39	0.09	3.30
AC 08	TCp	Bt2	10YR	2.6	55.9 (79.9)	17.3	1.3	0.31	0.08	2.79
AC 09	PVAal-1	Bt1	5YR	3.4	44.1 (63.1)	10.4	1.1	0.24	0.11	4.32
AC 09	PVAal-1	Bt2	5YR	2.5	30.1 (43.0)	9.7	0.8	0.32	0.08	4.40
AC 10	PVAal-2	Bt1	5YR	4.7	53.8 (76.9)	24.9	1.4	0.46	0.06	2.00
AC 10	PVAal-2	Bt2	5YR	3.2	56.6 (80.9)	20.9	1.2	0.37	0.06	2.27
AC 11	VXo-2	Biv1	10YR	8.5	42.0 (60.1)	16.6	2.7	0.40	0.16	3.07
AC 11	VXo-2	Biv2	10YR	5.8	58.7 (83.9)	–	–	–	–	3.48

Fes: Fe extracted by H<sub>2</sub>SO<sub>4</sub>; (Fe<sub>2</sub>O<sub>3</sub>) Fe extracted by H<sub>2</sub>SO<sub>4</sub> expressed as oxide form; Fed: Fe extracted by DCB-Na; Feo: Fe extracted by ammonium oxalate;  $K_i = (1.7 \cdot \text{SiO}_2) / \text{Al}_2\text{O}_3$  (ANJOS et al., 2013); EKo: Espodosolo Humilúvico Órtico; PVal: Argissolo Vermelho Alítico; VXo: Vertissolo Háplico Órtico; TCp: Luvissole Cromico Pálico; PVAal: Argissolo Vermelho-Amarelo Alítico; VXo: Vertissolo Háplico Órtico (SANTOS et al., 2013).

Table 2 - Hematite/(hematite + goethite) ratio, Fe by Al substitution and mean coherence lengths (MCL<sub>hkl</sub>) of hematite (Hm) and goethite (Gt) in the soil samples.

Profile	Soil	Horizon	Hm/(Hm+Gt)	---Al <sup>3+</sup> substitution---		-----MCL <sub>hkl</sub> , nm-----			
				Hm -----mol mol <sup>-1</sup> -----	Gt	Hm(104)	Hm(110)	Gt(110)	Gt (111)
AC 01	EKo	Bhsx	–	–	–	–	–	–	–
AC 02	LAd	Bw2	0.17	–	0.29	–	–	42	86
AC 04	PVal-1	Bt1	0.38	0.05	0.19	103	29	50	158
AC 04	PVal-1	Bt2	0.35	0.05	0.19	154	–	42	–
AC 04	PVal-1	BCf1	0.52	0.05	0.18	77	39	59	136
AC 05	PVal-2	BA	0.45	0.05	0.14	103	59	74	158
AC 05	PVal-2	Bt1	0.47	0.05	0.14	–	39	27	95
AC 05	PVal-2	Bt2	0.47	0.05	0.18	77	47	50	119
AC 06	PVal-3	BA	0.49	0.08	0.16	69	–	42	119
AC 06	PVal-3	Bt1	0.37	0.08	0.14	123	47	37	119
AC 06	PVal-3	Bt3	0.39	0.05	0.16	77	39	33	–
AC 07	VXo-1	Biv	0.00	–	0.06	–	–	–	–
AC 07	VXo-1	BCv2	0.00	–	0.05	–	–	–	158
AC 08	TCp	Bt1	0.19	0.05	0.11	–	–	33	158
AC 08	TCp	Bt2	0.21	0.05	0.13	103	–	33	136
AC 09	PVAal-1	Bt1	–	–	–	–	–	–	–
AC 09	PVAal-1	Bt2	–	–	–	–	–	–	–
AC 10	PVAal-2	Bt1	0.57	0.05	–	–	29	50	–
AC 10	PVAal-2	Bt2	0.60	0.08	0.15	88	47	59	–
AC 11	VXo-2	Biv1	0.28	–	0.28	88	–	27	–
AC 11	VXo-2	Biv2	–	–	0.16	–	–	37	–

EKo: Espodosolo Humilúvico Órtico; PVal: Argissolo Vermelho Alítico; VXo: Vertissolo Háplico Órtico; TCp: Luvisolo Cromico Pálico; PVAal: Argissolo Vermelho-Amarelo Alítico; VXo: Vertissolo Háplico Órtico (SANTOS et al., 2013).

plaques in pedogenic hematite (SCHWERTMANN & KÄMPF, 1985).

Overall, the studied soils, which developed from Cenozoic sedimentary rocks in the Solimões formation (Acre state, Brazil), have low contents in total Fe and Fe in pedogenic iron oxides. The substantial proportion of Fe in weathered minerals suggests that the soils are at an intermediate stage of weathering, where crystalline iron oxides prevail in the following sequence: goethite>hematite>maghemite. The conclusion that the soils are only moderately developed is supported by the increased frequency of goethite and hematite

with medium degrees of Al<sup>3+</sup> substitution. Finally, goethite is present mainly in isodimensional forms and hematite as flat plaques.

#### DECLARATION OF CONFLICTING OF INTERESTS

The authors have no conflict of interest to declare.

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